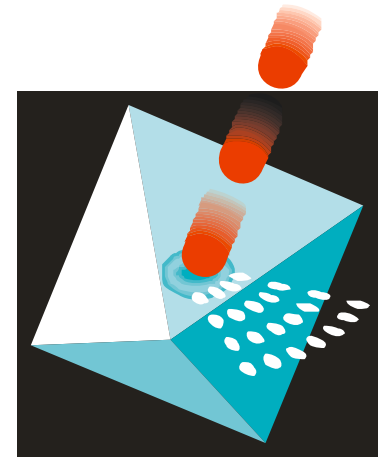


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# U-Th-Pb Geochronology Short Course

## Part 1: Basics



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*Funding from NSF (Instrumentation & Facilities) & University of Arizona*  
*Help from GV Instruments, Nu Plasma Instruments, New Wave Research*  
*& Photon Machines*

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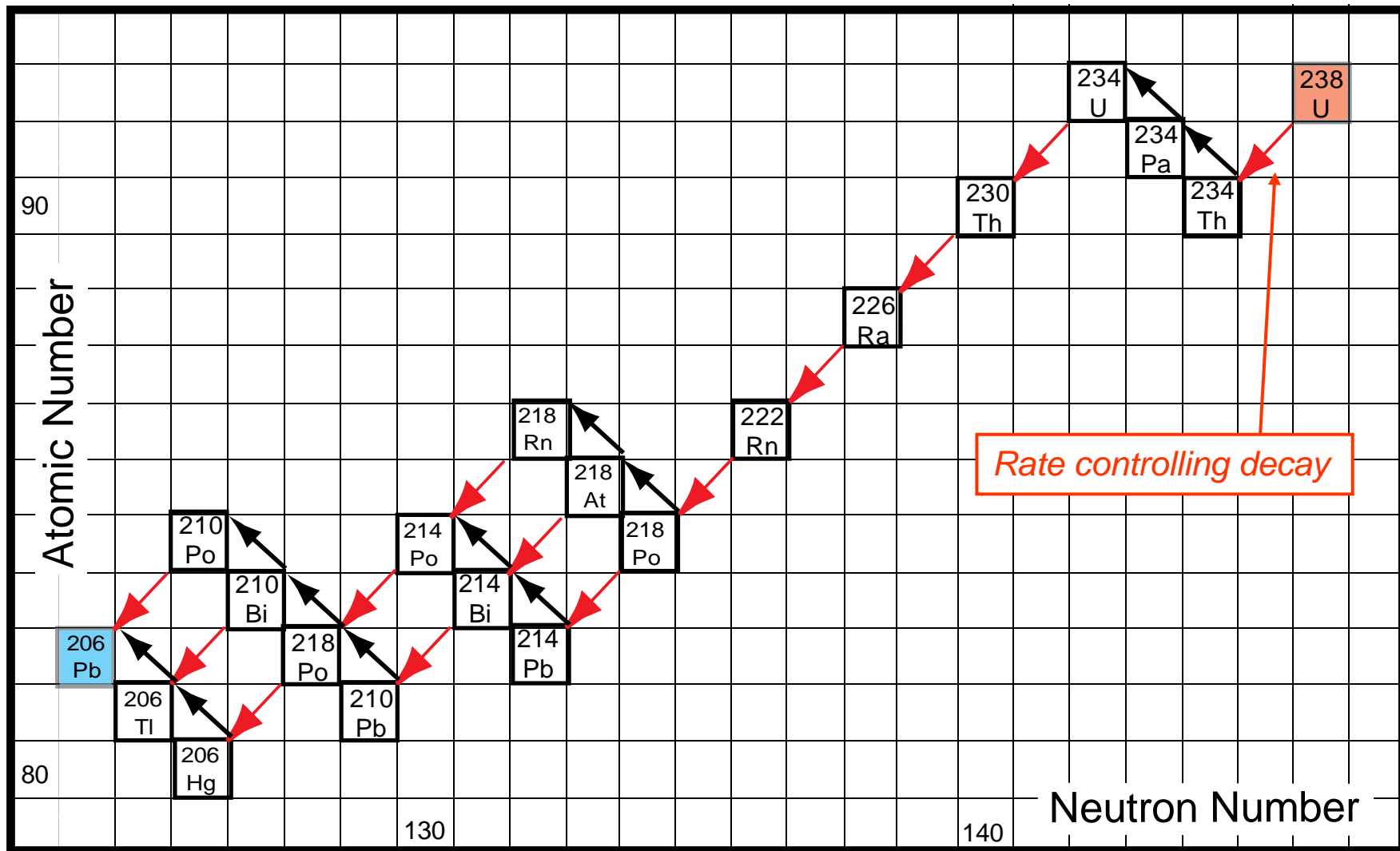
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# Outline

1. Basics of U-Th-Pb decay system
  2. Measurement methods
  3. Mineral systems & applications
  4. Complexities for zircon
  5. Detrital geochronology (plots & analyses)
  6. Data analysis
  7. Future opportunities
-

# Decay of $^{238}\text{U}$ to $^{206}\text{Pb}$

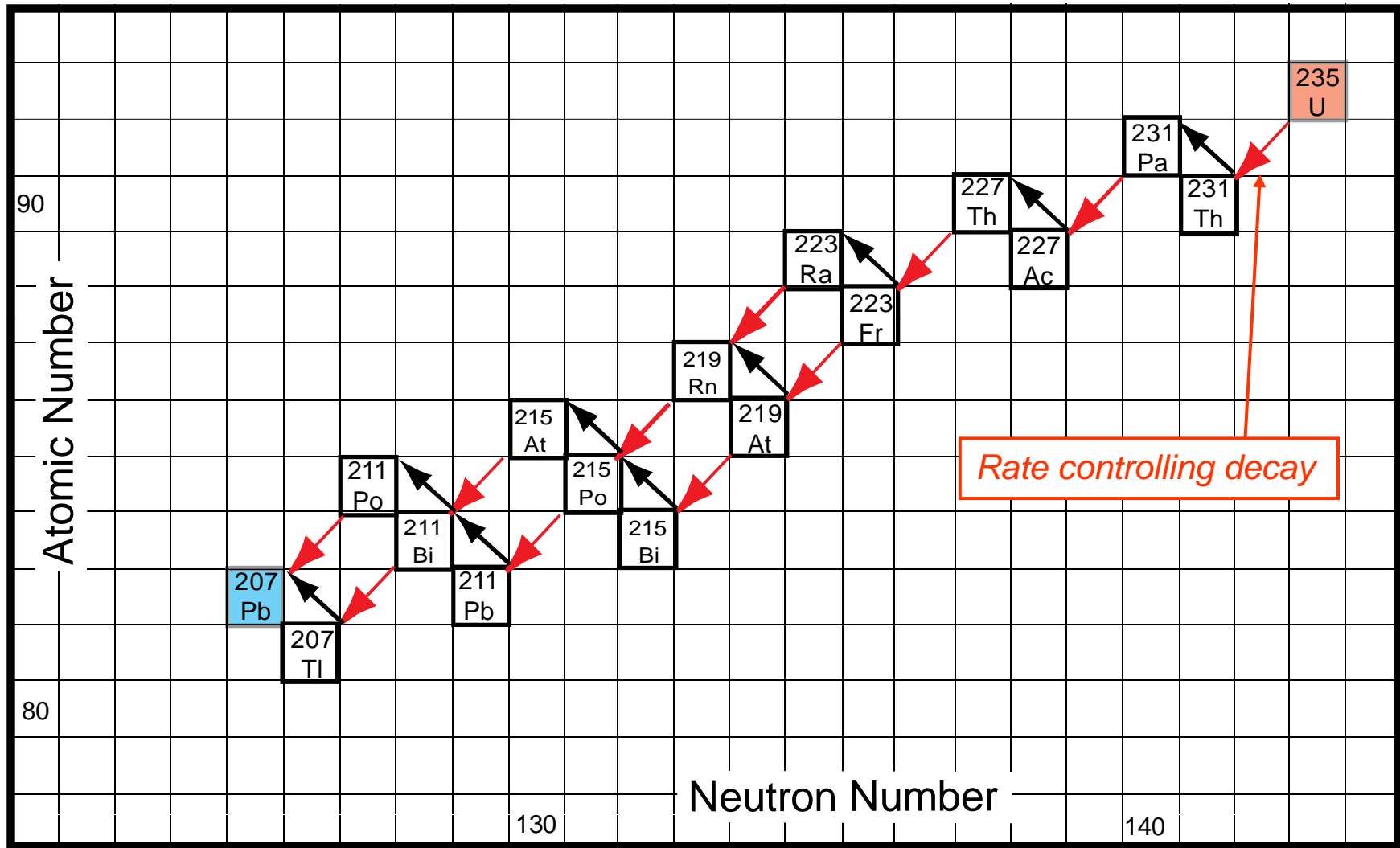
( $T_{1/2} = 4.468 \text{ Ga}$ )



*Possible complications from  $^{230}\text{Th}$ .....*

# Decay of $^{235}\text{U}$ to $^{207}\text{Pb}$

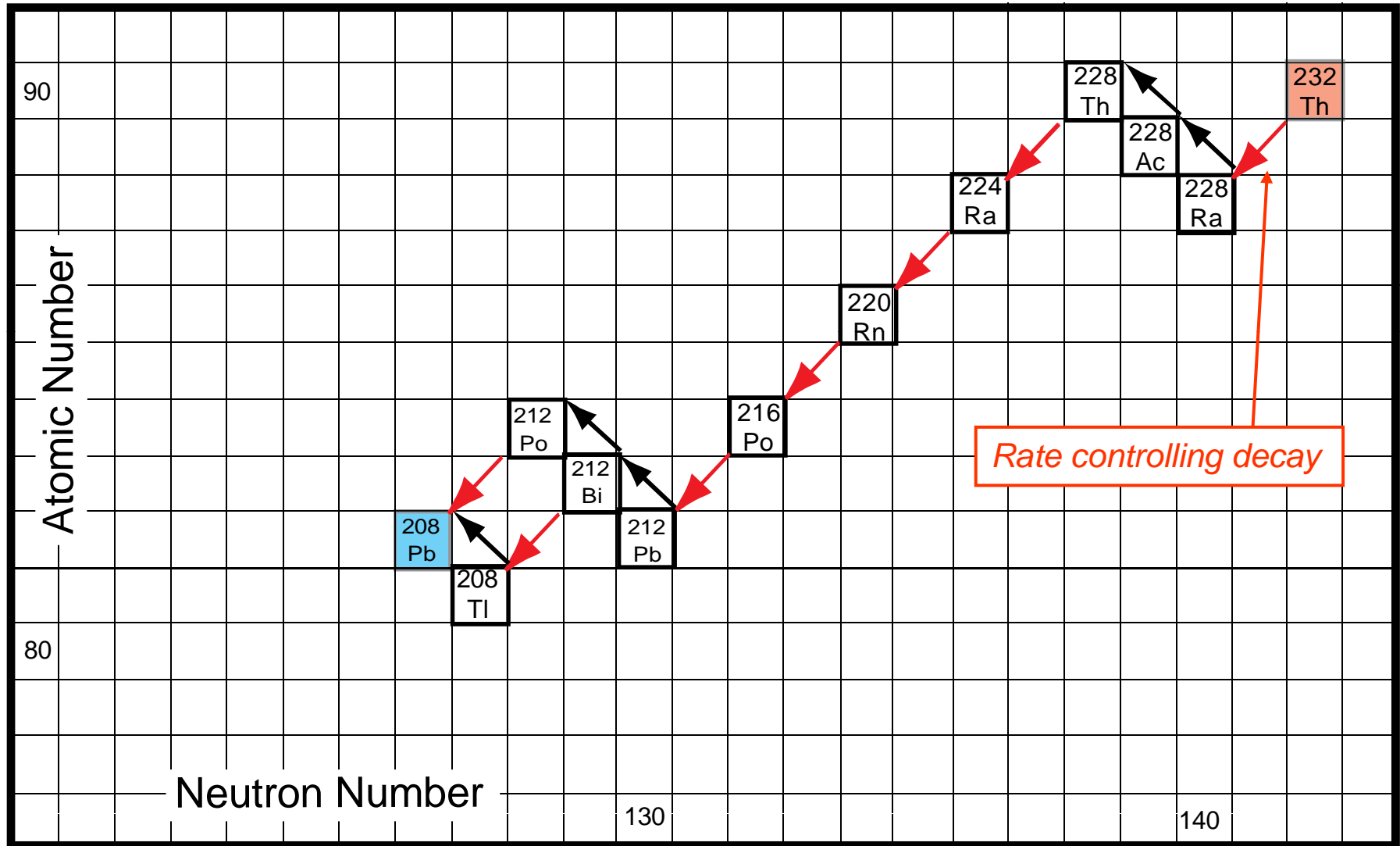
( $T_{1/2} = 0.704 \text{ Ga}$ )



Possible complications from  $^{231}\text{Protactinium}$

# Decay of $^{232}\text{Th}$ to $^{208}\text{Pb}$

( $T_{1/2} = 14.01 \text{ Ga}$ )



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## U-Th-Pb decay

$dN/dt =$  proportional to  $N$  (process of radioactive decay)

$$-dN/dt = \lambda N \rightarrow \text{(integrate both sides)} \rightarrow -\ln N = \lambda t + c$$

because  $N = N_i$  when  $t = 0$ ,  $c = -\ln N_i$

$$-\ln N = \lambda t - \ln N_i \rightarrow N = N_i e^{-\lambda t}$$

$$D^* = N_i - N \rightarrow N_i = D^* + N$$

$$D^* = N (e^{\lambda t} - 1) \rightarrow \quad \quad \quad \mathbf{D^*/N = e^{\lambda t} - 1 \text{ (decay equation)}}$$

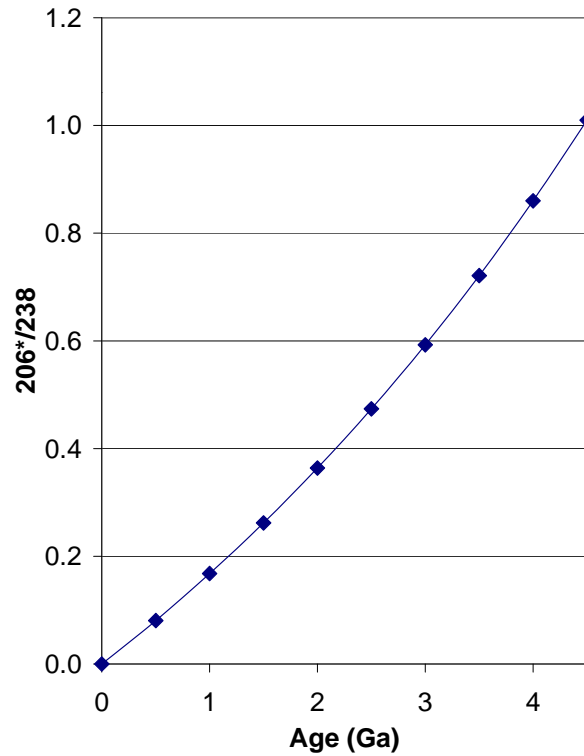
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# D\*/N as function of age:



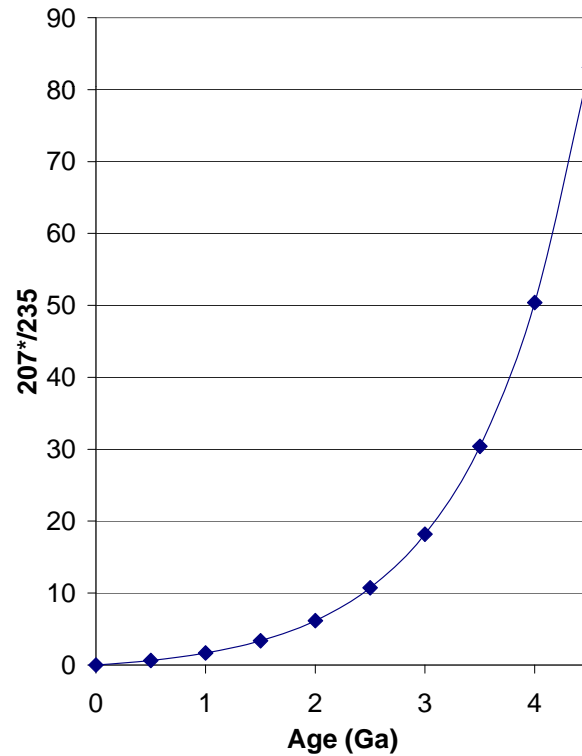
$T_{1/2} = 4.468 \text{ Ga}$

$^{206}\text{Pb}^*/^{238}\text{U} = e^{\lambda_1 t} - 1$



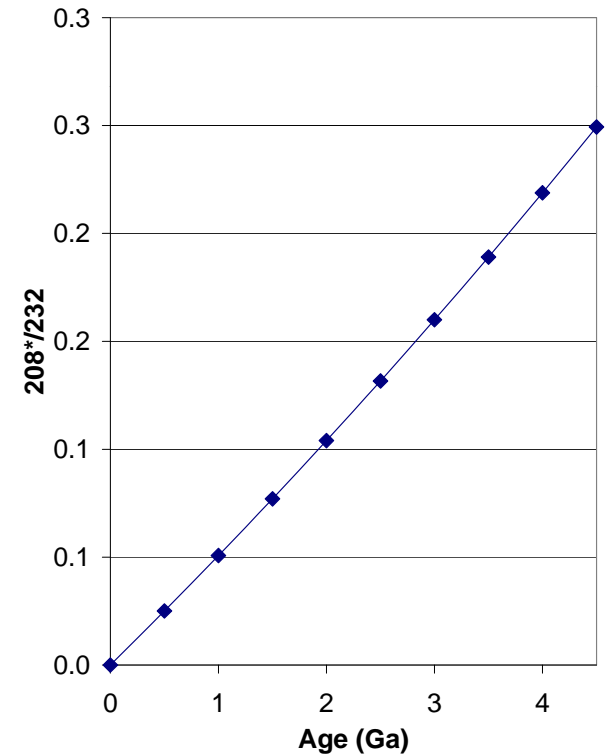
$T_{1/2} = 0.704 \text{ Ga}$

$^{207}\text{Pb}^*/^{235}\text{U} = e^{\lambda_2 t} - 1$



$T_{1/2} = 14.01 \text{ Ga}$

$^{208}\text{Pb}^*/^{232}\text{Th} = e^{\lambda_3 t} - 1$



\* = radiogenic

# U-Th-Pb age equations

$$D^*/N = e^{\lambda t} - 1 \text{ (decay equation)}$$

solve for t

$$t = \ln (D^*/N + 1) / \lambda \text{ (age equation)}$$

$$\text{Age}_{(206/238)} = \ln (206^*/238 + 1) / {}^{238}\lambda$$

$$\text{Age}_{(207/235)} = \ln (207^*/235 + 1) / {}^{235}\lambda$$

$$\text{Age}_{(208/232)} = \ln (208^*/232 + 1) / {}^{232}\lambda$$

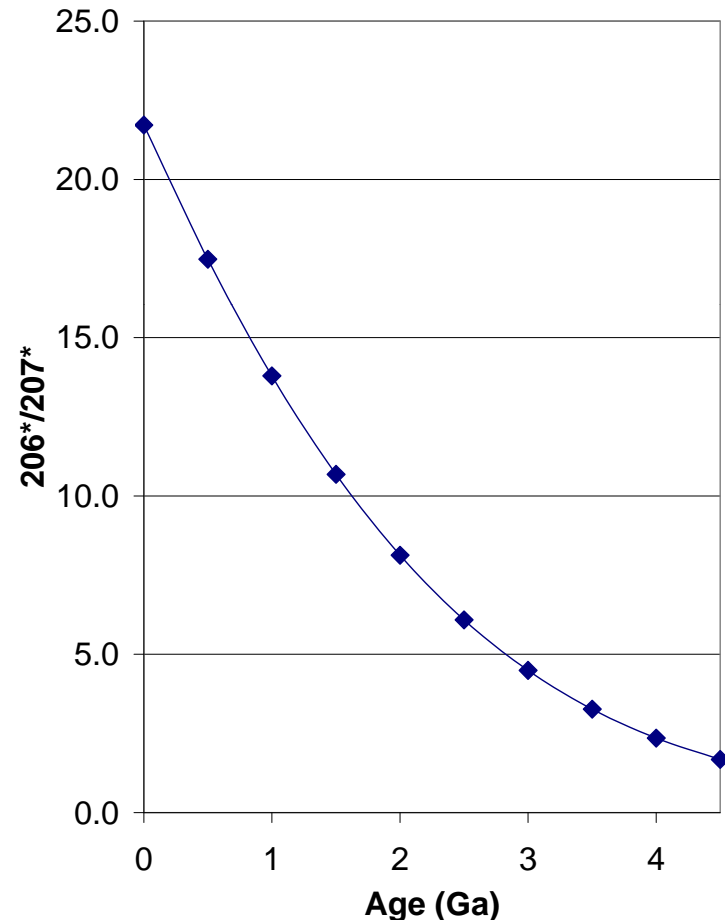
Can also determine age from 206/207  
because  $238/235 = 137.88$  in nearly all rocks:

$${}^{206}\text{Pb}^* = {}^{238}\text{U} (e^{\lambda_1 t} - 1)$$

$${}^{207}\text{Pb}^* = {}^{235}\text{U} (e^{\lambda_2 t} - 1)$$

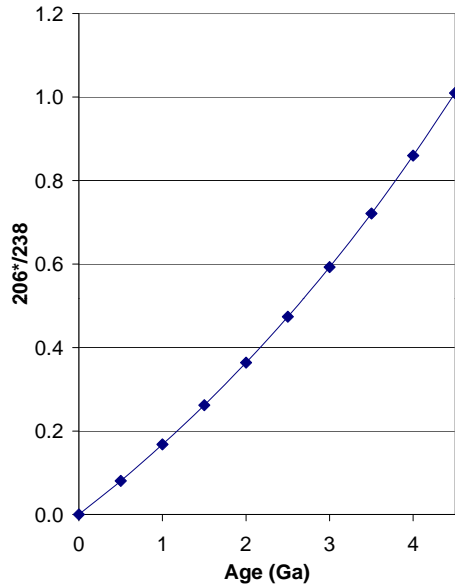
$$\begin{aligned} {}^{206}\text{Pb}^* / {}^{207}\text{Pb}^* &= [{}^{238}\text{U} (e^{\lambda_1 t} - 1)] / [{}^{235}\text{U} (e^{\lambda_2 t} - 1)] \\ &= 137.88 [(e^{\lambda_1 t} - 1) / (e^{\lambda_2 t} - 1)] \end{aligned}$$

*(have to iterate or use a table.....)*

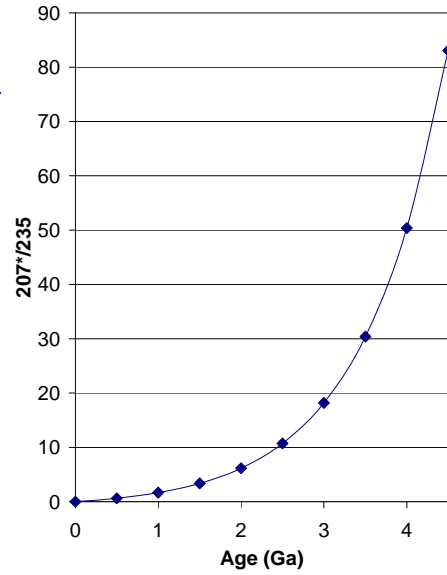


# U-Th-Pb chronometers:

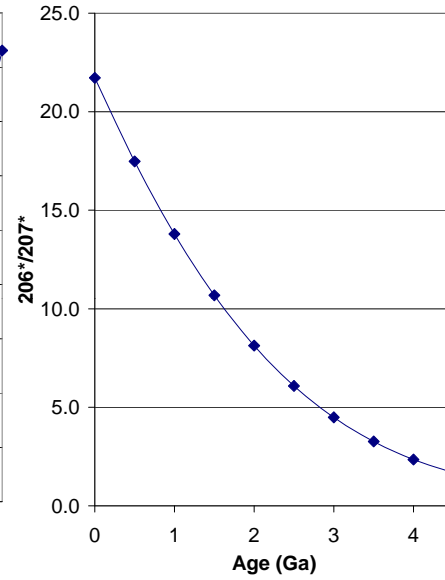
206/238



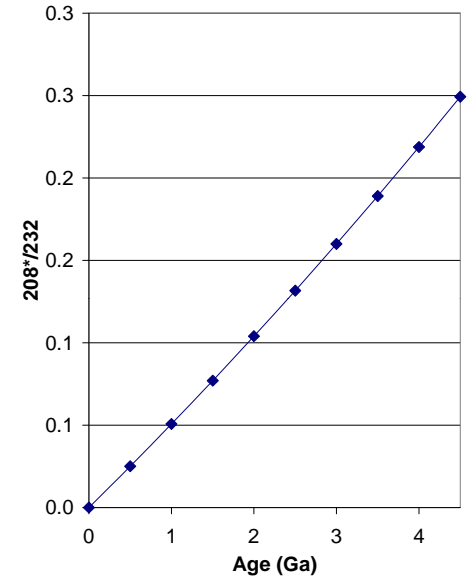
207/235



206/207



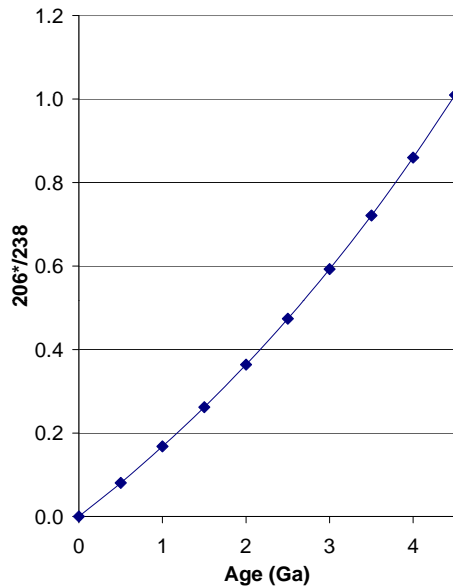
208/232



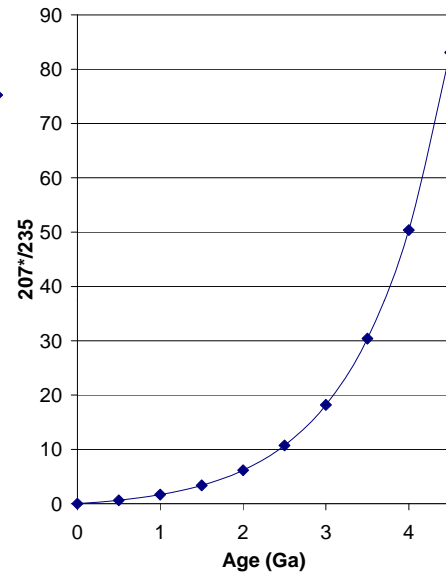
# U-Th-Pb chronometers:

Low Th in most minerals  
→ Tough to measure!

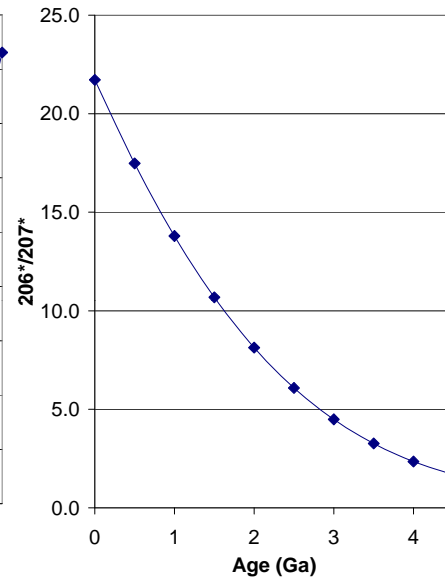
206/238



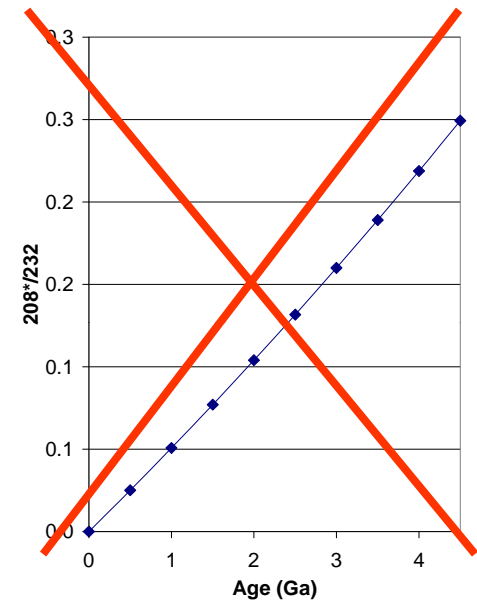
207/235



206/207



208/232



# What about Pb in crystal when it formed?

Measured Pb = radiogenic Pb + common Pb (present when crystal formed):

$$^{204}\text{M} = ^{204}\text{C}$$

$$^{206}\text{M} = ^{206}\text{C} + ^{206}\text{R}$$

$$^{207}\text{M} = ^{207}\text{C} + ^{207}\text{R}$$

$$^{208}\text{M} = ^{208}\text{C} + ^{208}\text{R}$$

$$\left(\frac{206}{204}\right)\text{M} = \left(\frac{206}{204}\right)\text{R} + \left(\frac{206}{204}\right)\text{C}$$

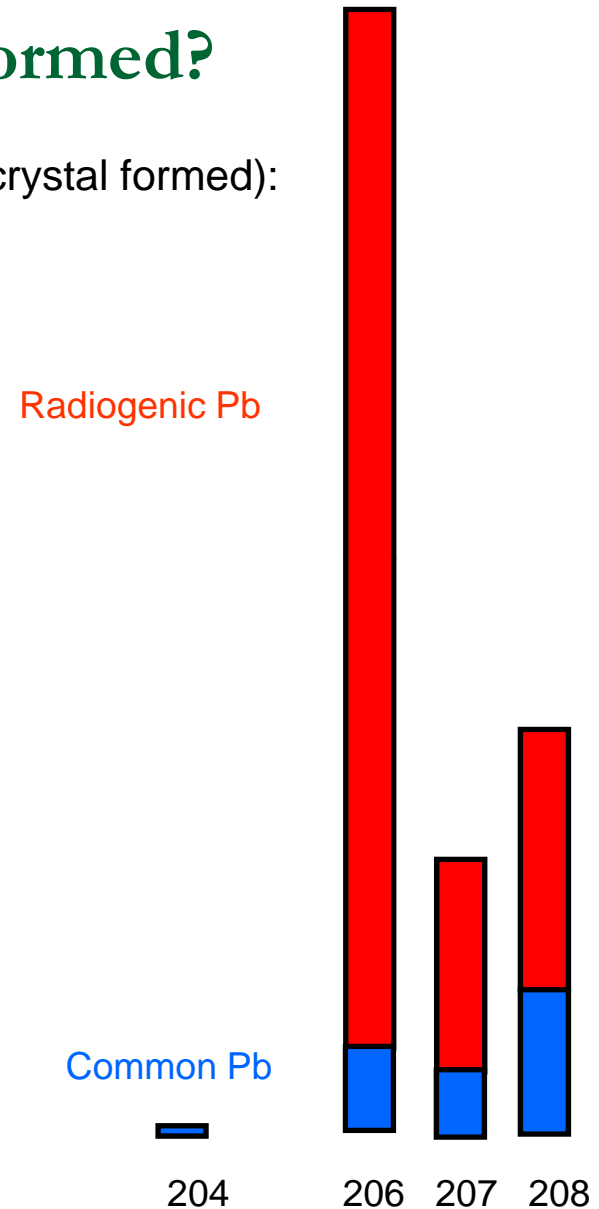
$$\rightarrow \left(\frac{206}{204}\right)\text{R} = \left(\frac{206}{204}\right)\text{M} - \left(\frac{206}{204}\right)\text{C}$$

$$\rightarrow 206^* = 204 \left[ \left(\frac{206}{204}\right)\text{M} - \left(\frac{206}{204}\right)\text{C} \right]$$

$$\rightarrow 207^* = 204 \left[ \left(\frac{207}{204}\right)\text{M} - \left(\frac{207}{204}\right)\text{C} \right]$$

$$\rightarrow 208^* = 204 \left[ \left(\frac{208}{204}\right)\text{M} - \left(\frac{208}{204}\right)\text{C} \right]$$

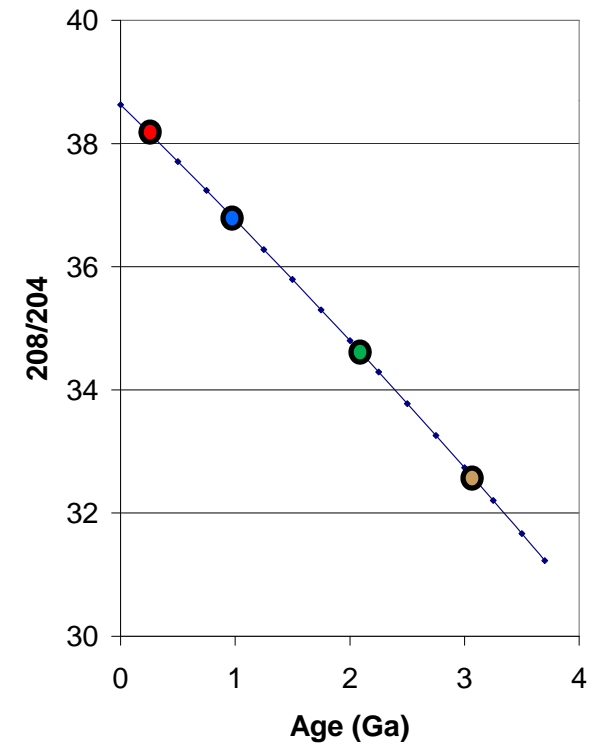
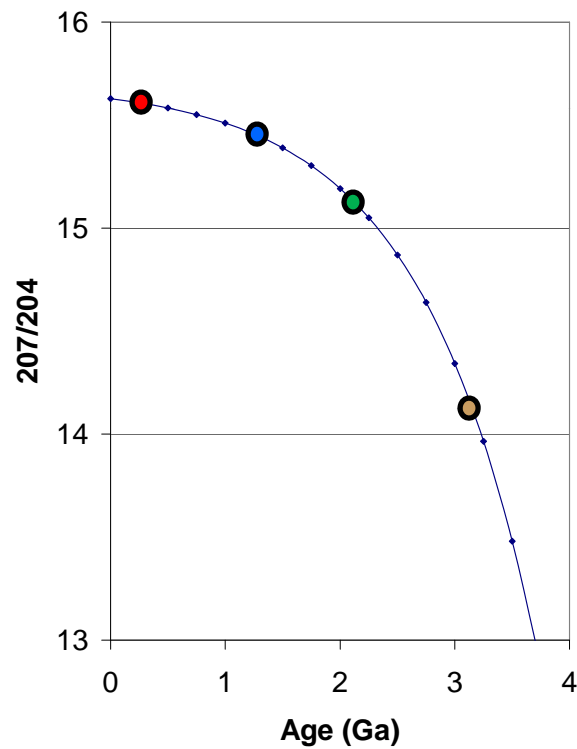
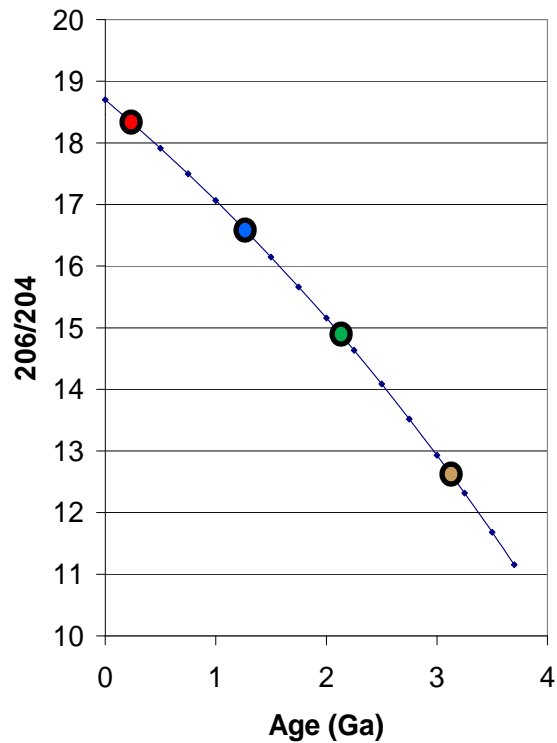
*How determine common Pb composition when  
Crystal formed??*



# Common Pb composition

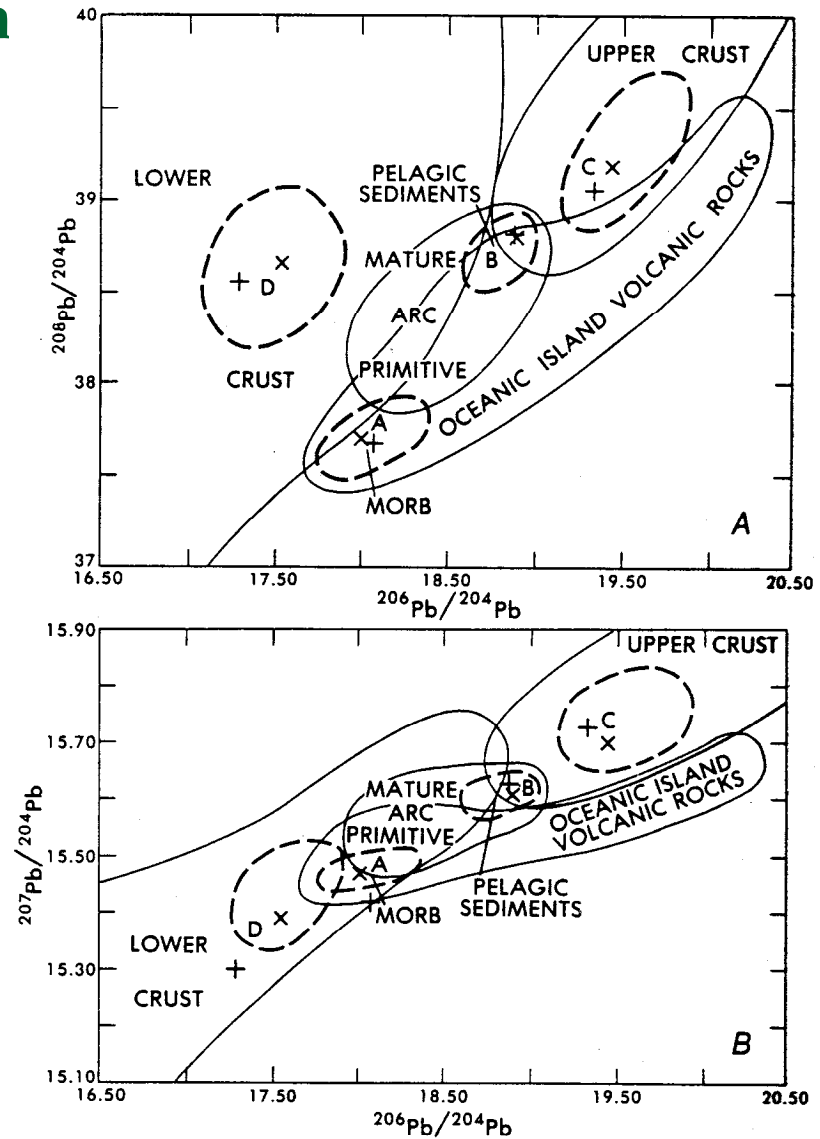
- Ore minerals (e.g. galena) have such high Pb contents that they do not change over time
- Measure Pb in ore minerals of various ages to determine evolution of Pb in the crust ●

= Stacey-Kramers (1975) model for evolution of crustal Pb



# Common Pb composition

But there is variation in modern common Pb composition:



# Common Pb composition

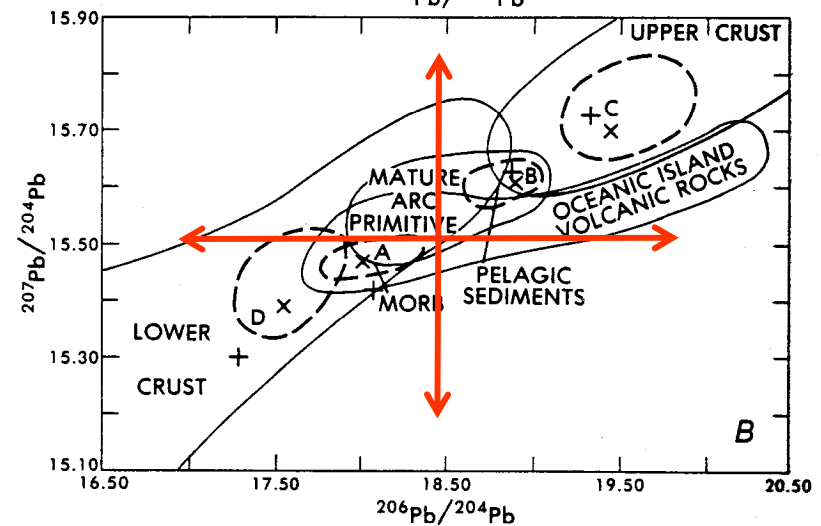
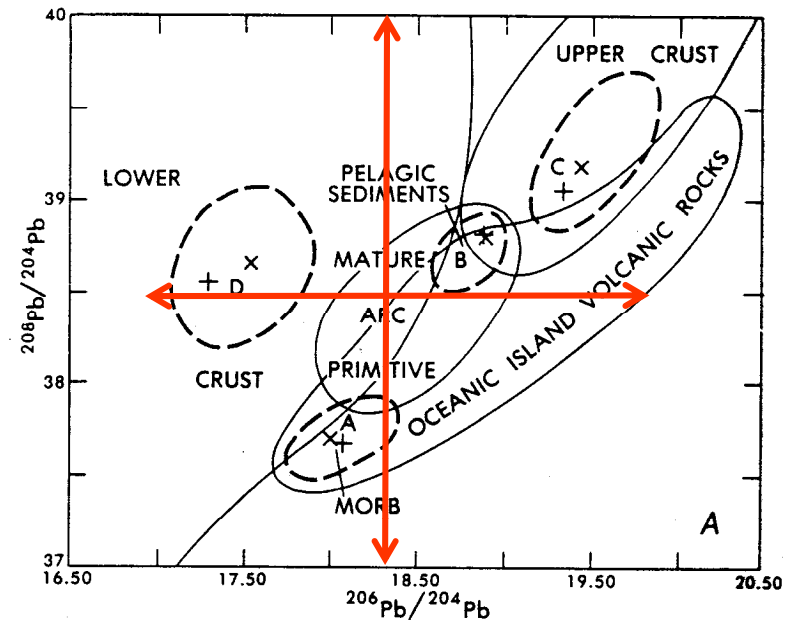
But there is variation in common Pb composition today:

So need to apply an uncertainty to Stacey-Kramers Pb compositions.....

$\pm 1.5$  to  $^{206}\text{Pb}/^{204}\text{Pb}$

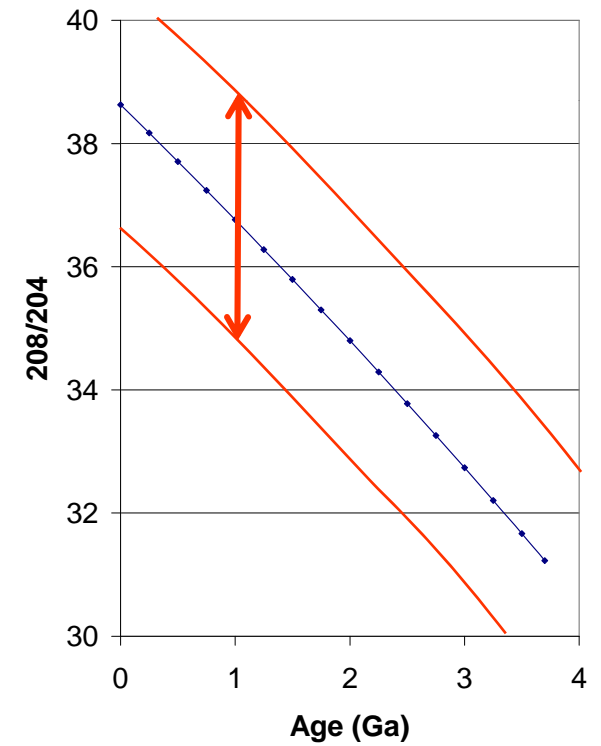
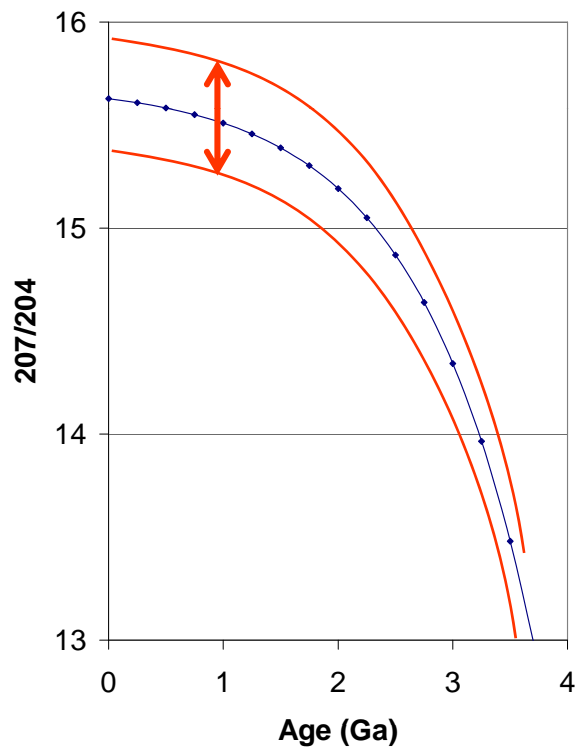
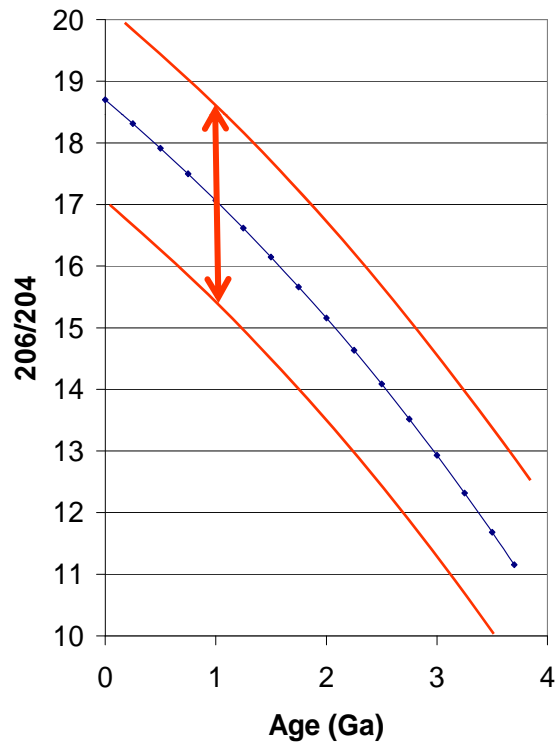
$\pm 0.3$  to  $^{207}\text{Pb}/^{204}\text{Pb}$

$\pm 1.5$  to  $^{208}\text{Pb}/^{204}\text{Pb}$



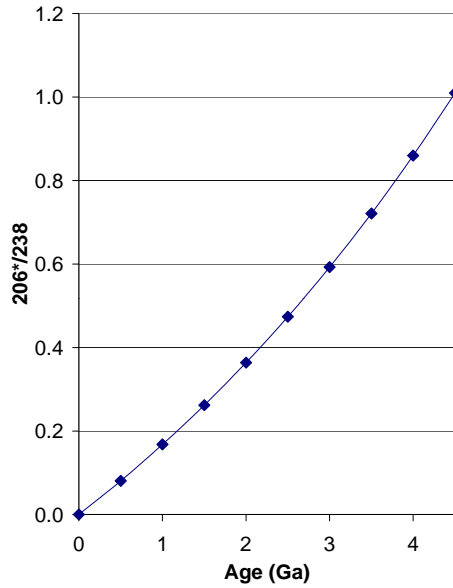
# Common Pb correction

→ Apply Stacey-Kramers (1975) composition with uncertainties  
(for age of crystallization = slightly circular.....)

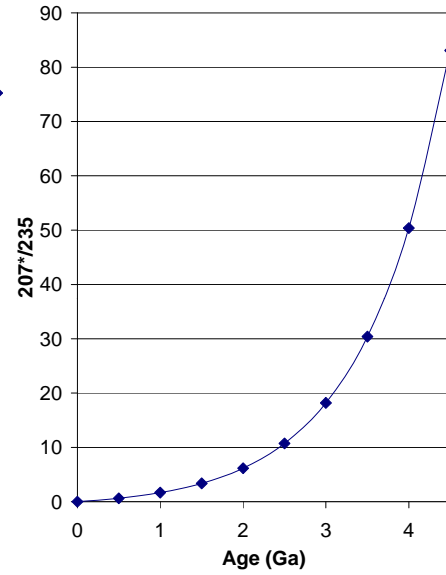


# Plot U-Pb chronometers on a single diagram:

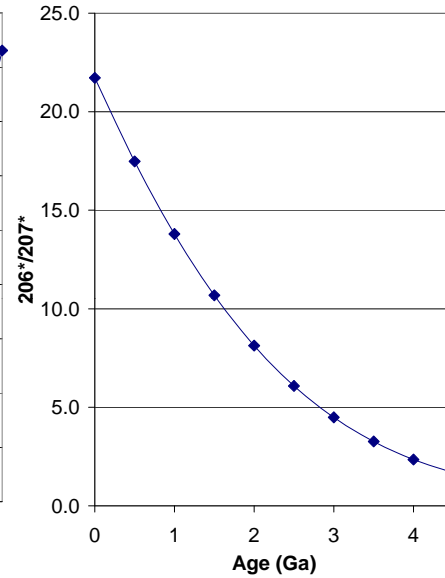
206/238



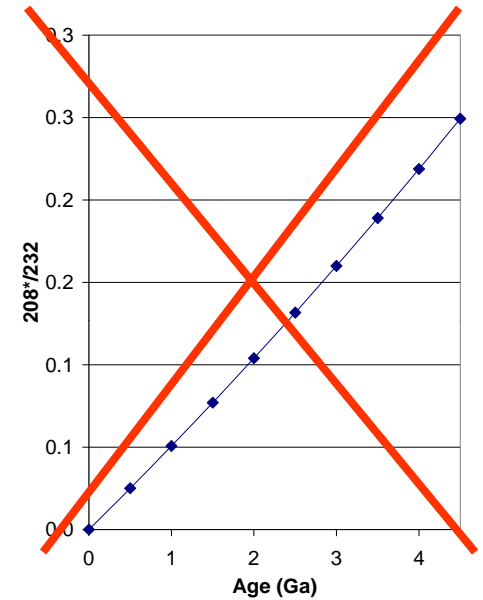
207/235



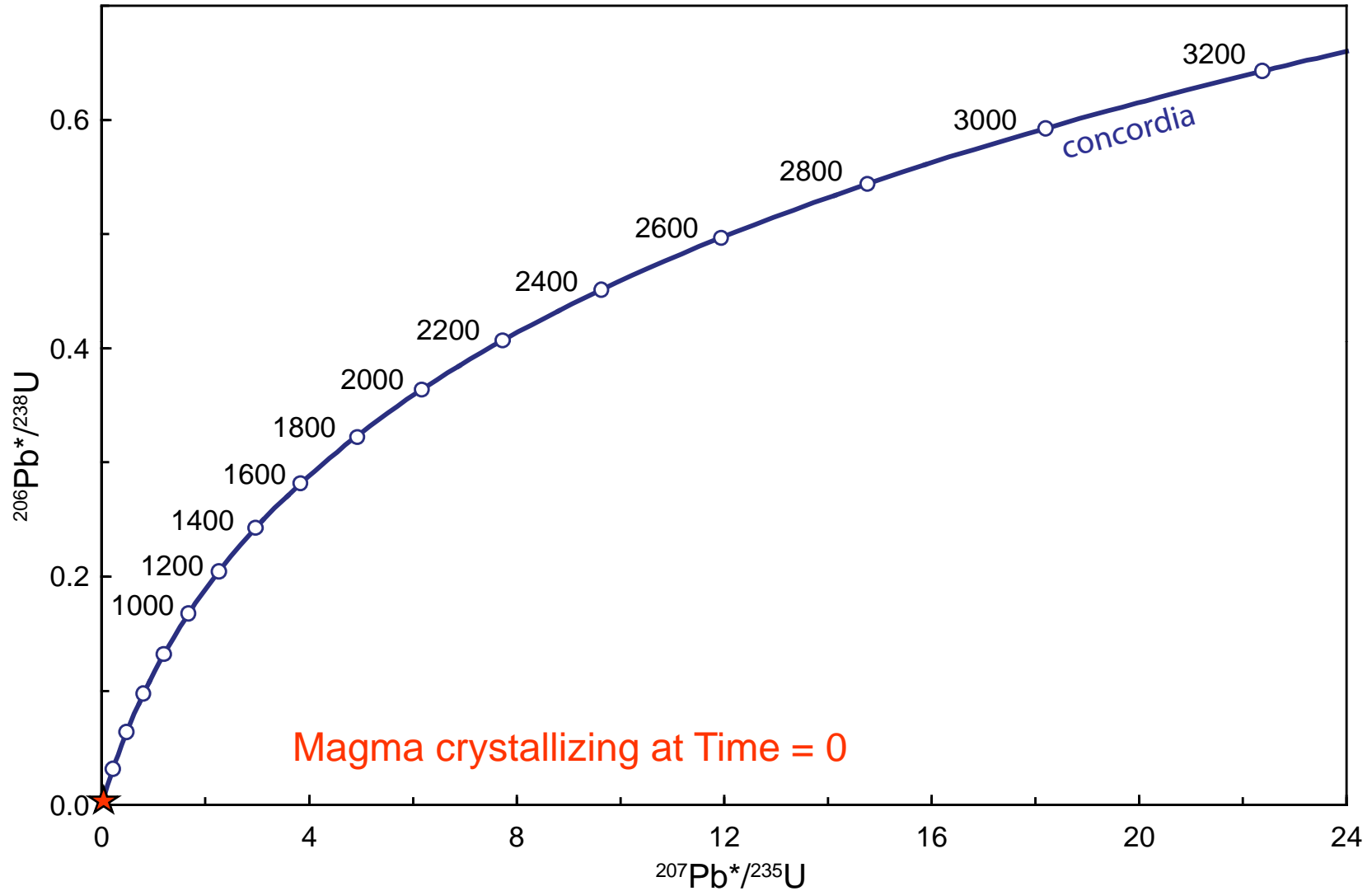
206/207



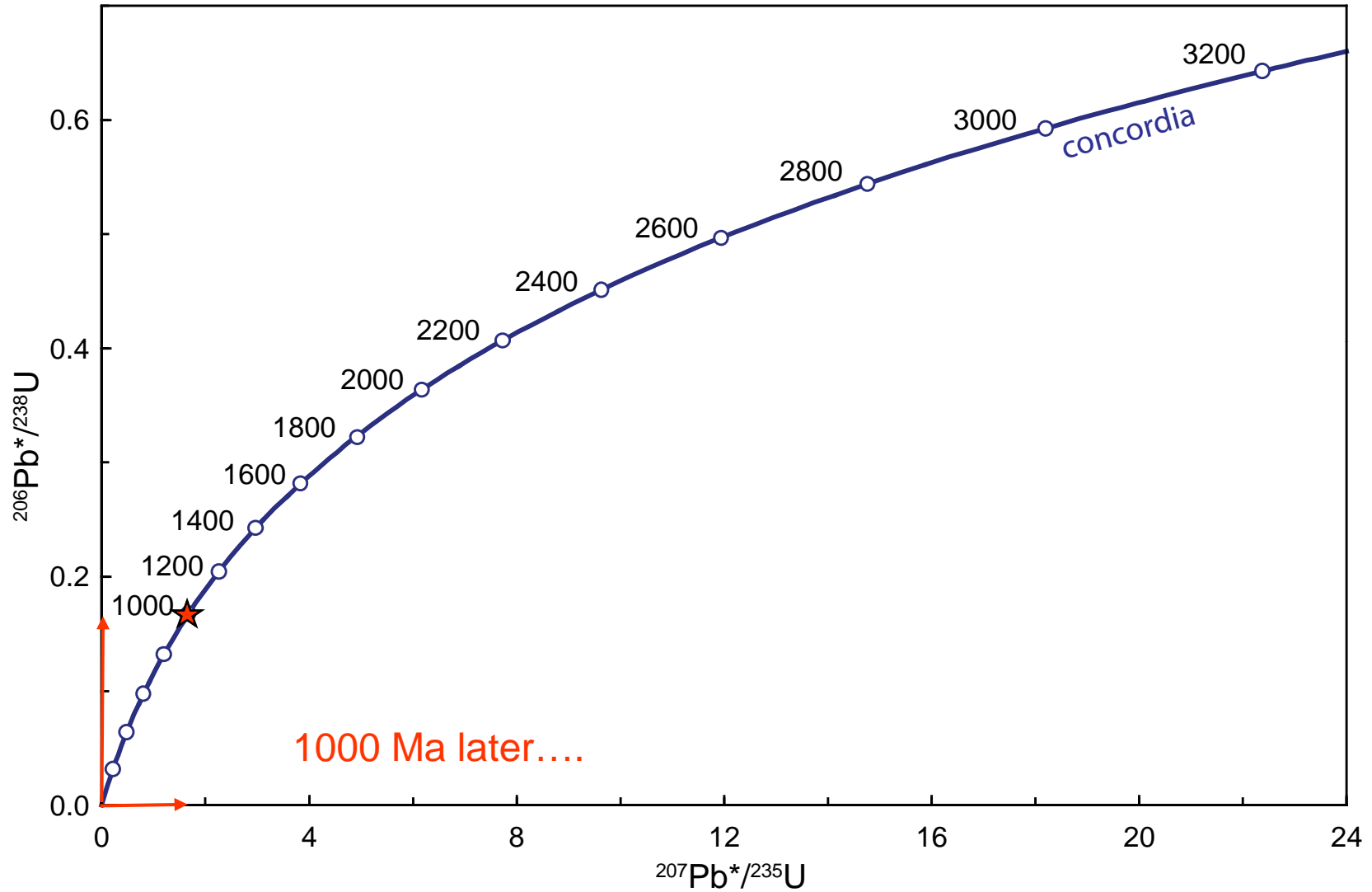
208/232



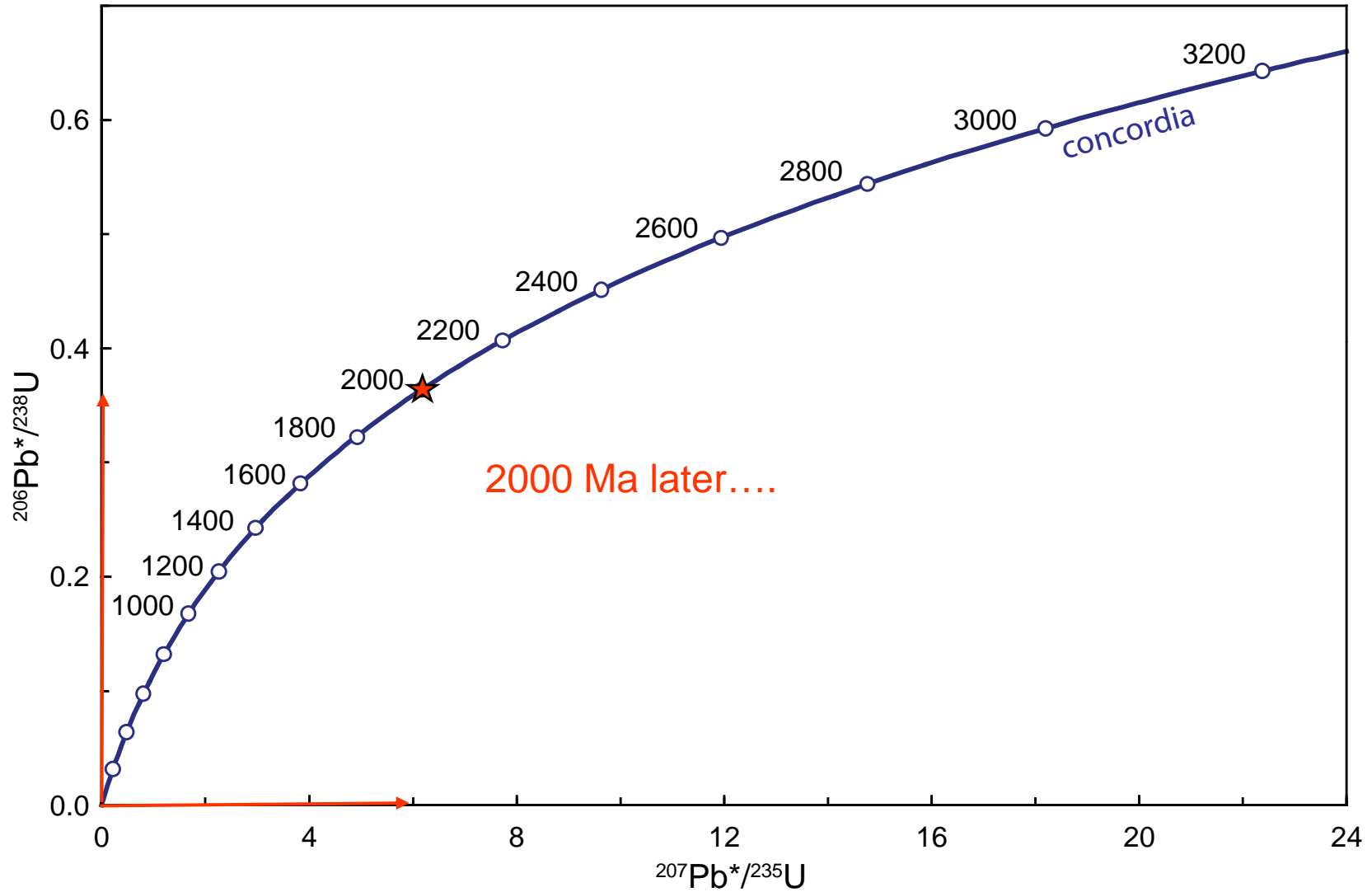
# U-Pb concordia diagram (*Wetherill, 1956*)



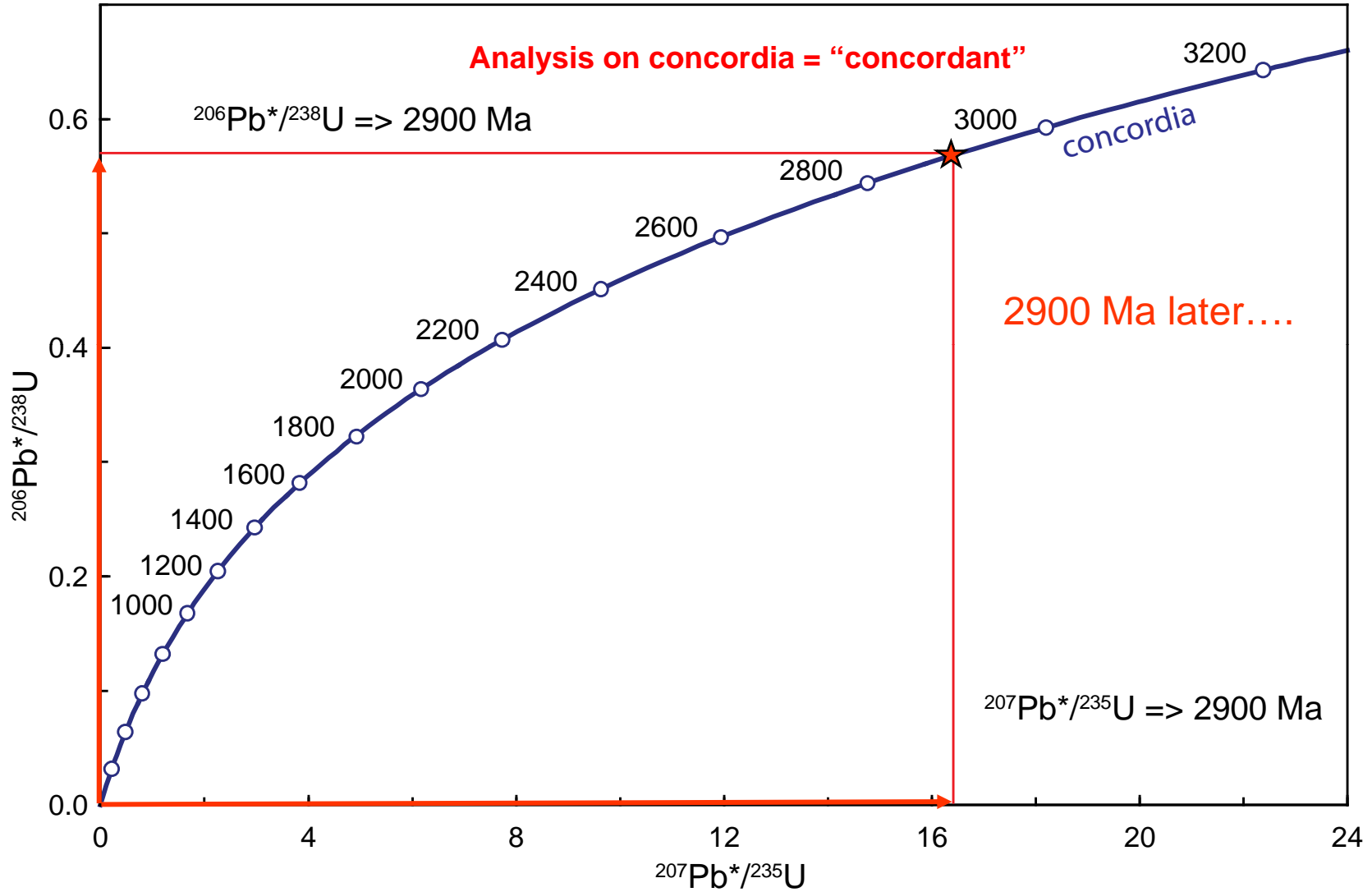
# U-Pb concordia diagram



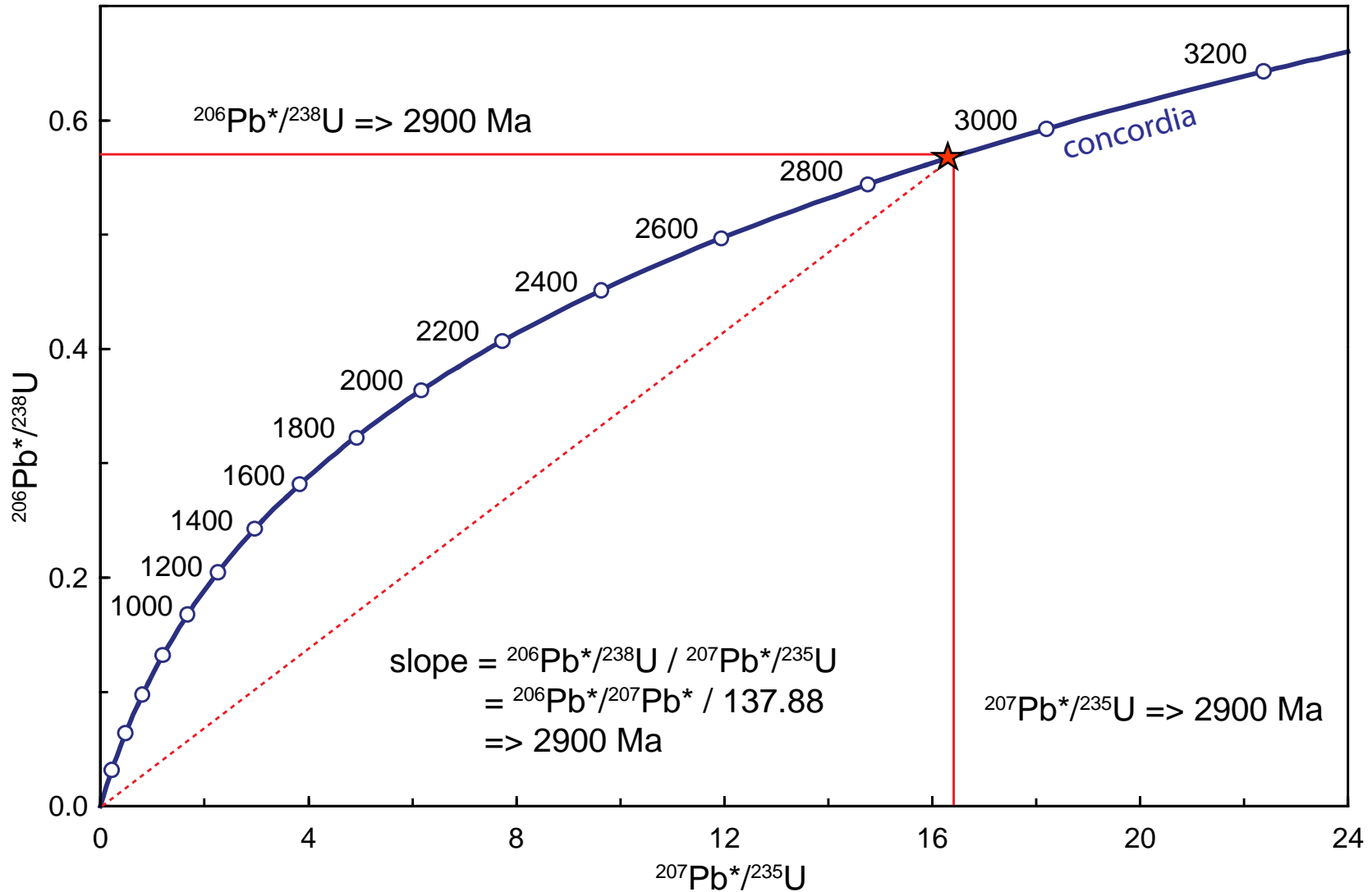
# U-Pb concordia diagram



# U-Pb concordia diagram

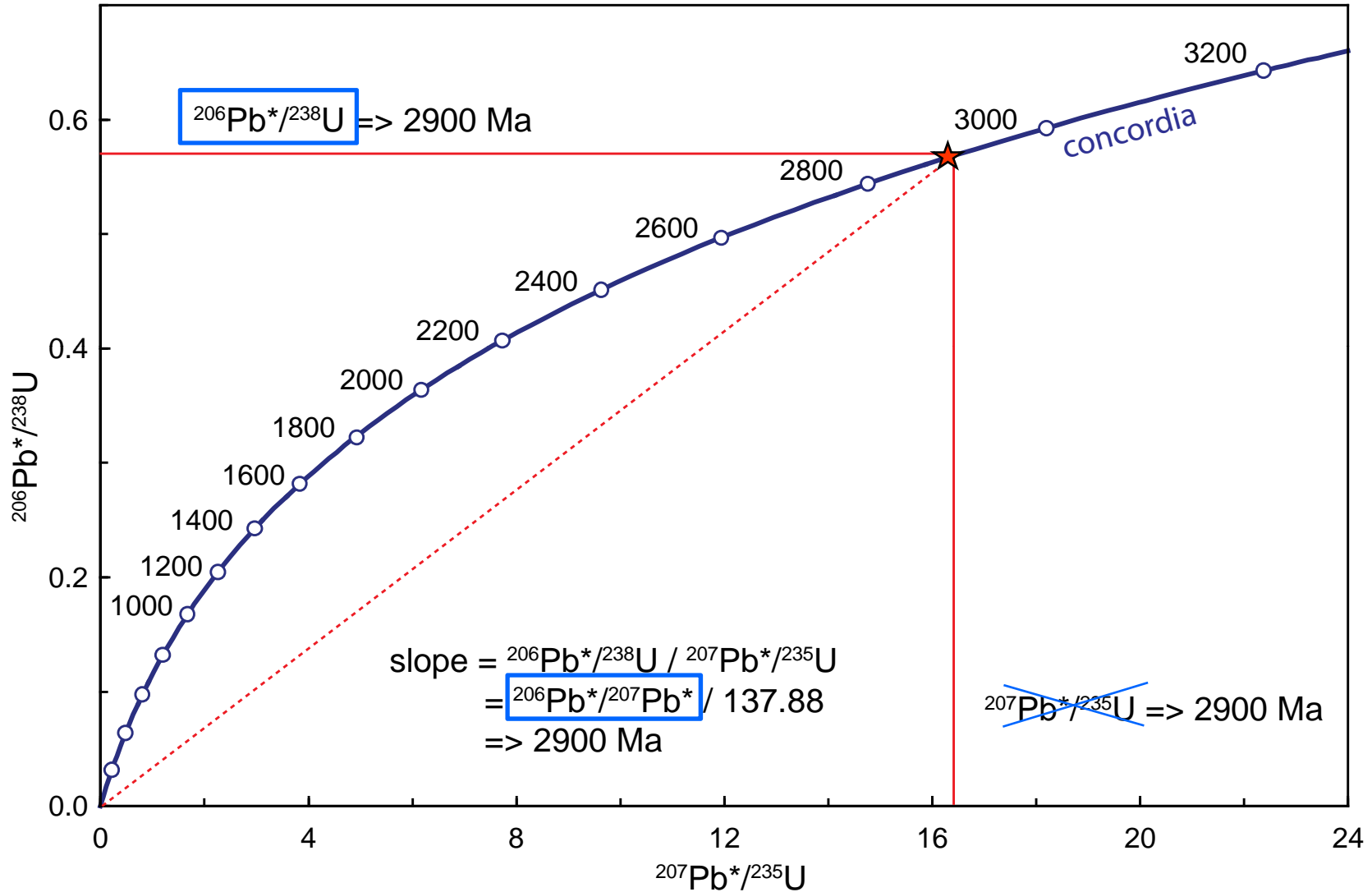


# U-Pb concordia diagram → 3 ages!



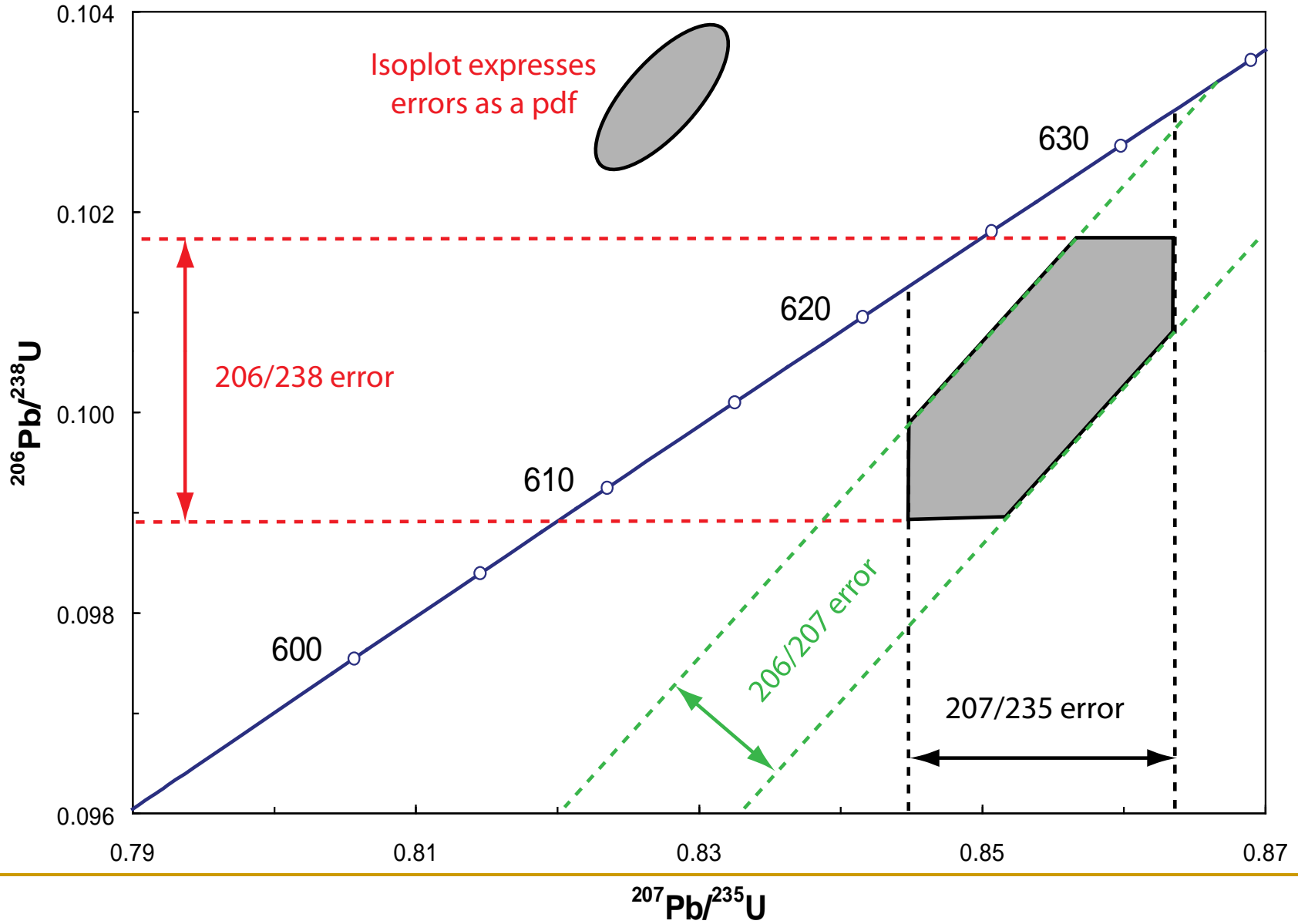
# U-Pb concordia diagram:

$^{206}\text{Pb}^*/^{238}\text{U}$  vs  $^{206}\text{Pb}^*/^{207}\text{Pb}^*$

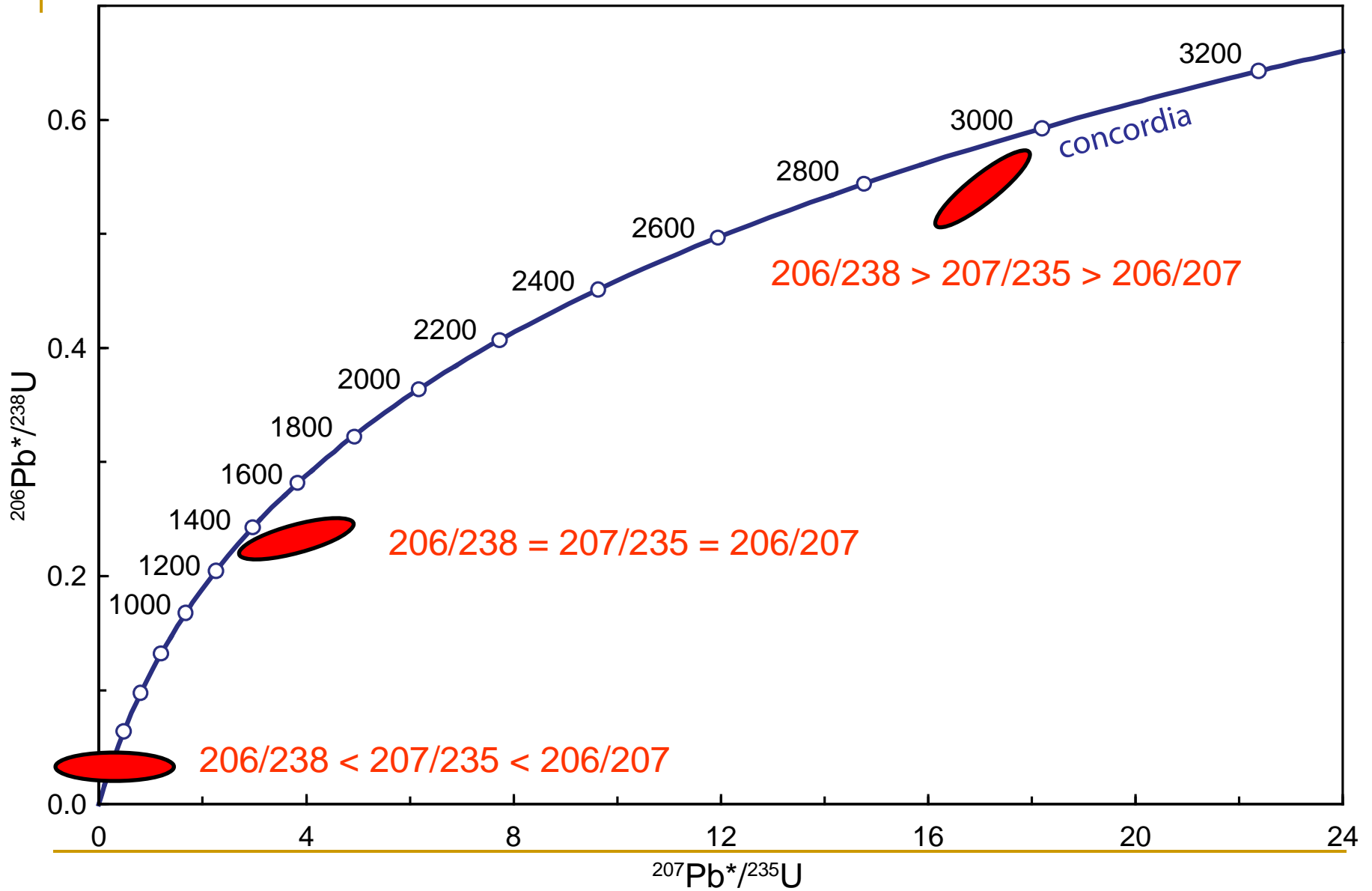


Don't measure  $^{207}\text{Pb}^*/^{235}\text{U}$  (=  $^{206}\text{Pb}^*/^{238}\text{U} / ^{206}\text{Pb}^*/^{207}\text{Pb}^* / 137.88$ )

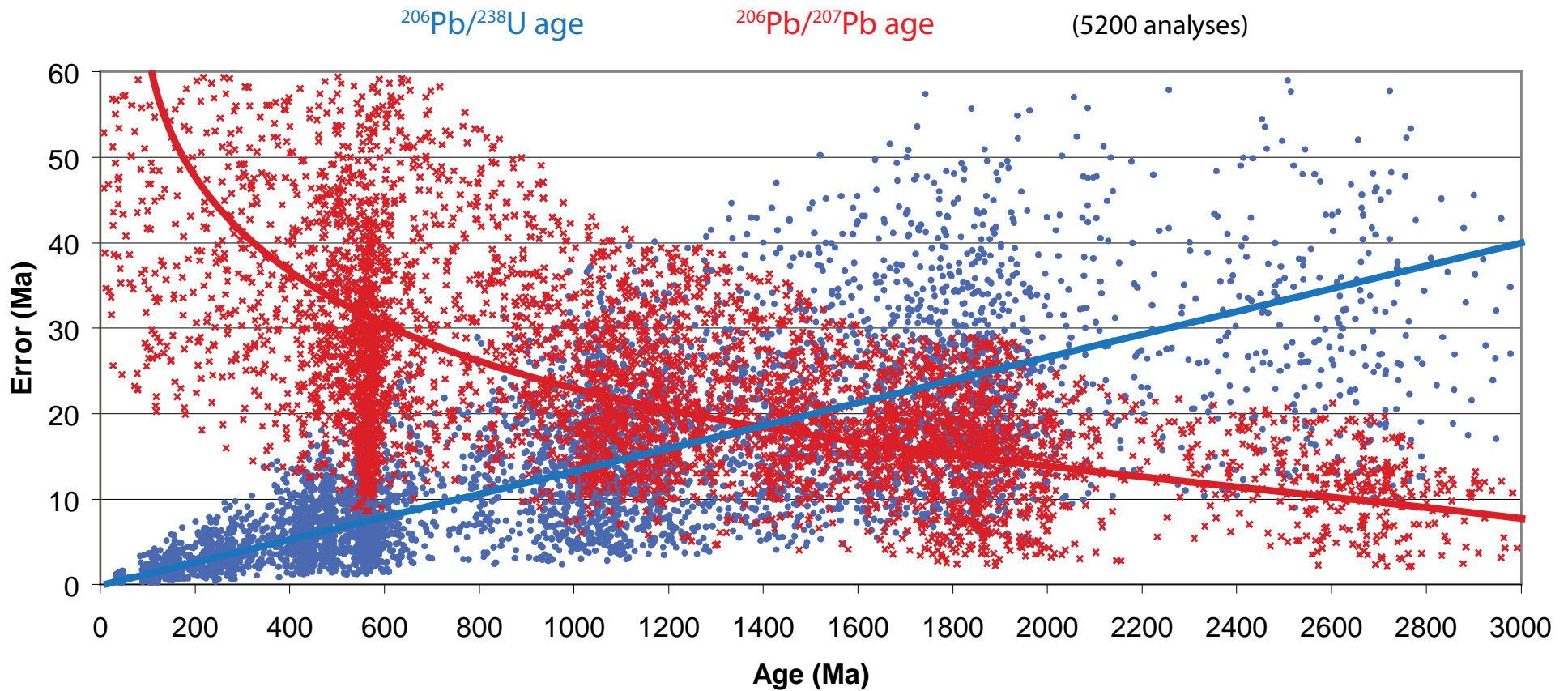
# Errors on U-Pb concordia diagram



# Errors on U-Pb concordia diagram



# Which age to use??



$^{206}\text{Pb}^*/^{238}\text{U}$  for ages < ~1.4 Ga

$^{206}\text{Pb}^*/^{207}\text{Pb}^*$  for ages > ~1.4 Ga

*Precision only* → *cutoff = 1.4 Ga*